An Exposition of Cessi's Ocean Box Model

Kate Meyer

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Outline

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Understanding ocean circulation Modeling resilience to transient parameter changes

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Cessi's motivation: understanding ocean circulation



http://earthobservatory.nasa.gov/Features/Paleoclimatology_Evidence/paleoclimatology_evidence_2.php

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Cessi's motivation: understanding ocean circulation

In models, competing influences of temperature and salinity give rise to alternative stable flow states:



e.g. Stommel (1961), Welander (1986), Manabe and Stouffer (1988)

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Cessi's motivation: understanding ocean circulation

The Younger Dryas

A period of abnormally cold temperatures that occurred between 12,900 and 11,500 years BP during the last deglaciation



Broecker, Wallace S. (2006). "Was the Younger Dryas Triggered by a Flood?". Science 312 (5777): 11461148

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Cessi's motivation: understanding ocean circulation altered salinity forcing



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My motivation

Modeling resilience to

transient parameter changes



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The box model: setup



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The box model: setup



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The box model: dimension reduction

$$\begin{aligned} T_1' &= -\frac{1}{t_r} \left(T_1 - \frac{\theta}{2} \right) - \frac{1}{2} Q(\Delta \rho) (T_1 - T_2) \\ T_2' &= -\frac{1}{t_r} \left(T_1 + \frac{\theta}{2} \right) - \frac{1}{2} Q(\Delta \rho) (T_2 - T_1) \\ \end{array} \qquad \qquad S_1' = -\frac{F(t)}{2H} S_0 - \frac{1}{2} Q(\Delta \rho) (S_1 - S_2) \\ S_2' &= -\frac{F(t)}{2H} S_0 - \frac{1}{2} Q(\Delta \rho) (S_2 - S_1) \end{aligned}$$

$$\Delta T \equiv T_1 - T_2; \qquad \Delta S \equiv S_1 - S_2$$

$$\Delta T' = -\frac{1}{t_r} (\Delta T - \theta) - Q(\Delta \rho) \Delta T$$
$$\Delta S' = \frac{F(t)}{H} S_0 - Q(\Delta \rho) \Delta S$$

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The box model: dimension reduction

$$\Delta T' = -\frac{1}{t_r} (\Delta T - \theta) - Q(\Delta \rho) \Delta T$$

$$\Delta S' = \frac{F(t)}{H} S_0 - Q(\Delta \rho) \Delta S$$

$$\begin{bmatrix} \Delta \rho = \alpha_S (S_1 - S_2) - \alpha_T (T_1 - T_2) \\ Q(\Delta \rho) = \frac{1}{t_d} + \frac{q(\Delta \rho)^2}{V} \end{bmatrix}$$

$$x \equiv \frac{\Delta T}{\theta}, \quad y \equiv \frac{\alpha_S \Delta S}{\alpha_T \theta}, \quad t \equiv t_d t'$$

$$\alpha \equiv \frac{t_d}{t_r}, \quad p(t) \equiv \frac{\alpha_S S_0 t_d}{\alpha_T \theta H} F(t), \quad \mu^2 \equiv \frac{q t_d (\alpha_T \theta)^2}{V} \end{bmatrix}$$

$$x' = -\alpha(x-1) - x \left[1 + \mu^2(x-y)^2\right]$$

$$y' = p(t) - y \left[1 + \mu^2(x-y)^2\right]$$

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The box model: dimension reduction

$$x' = -\alpha(x-1) - x \left[1 + \mu^2 (x-y)^2\right]$$

$$y' = p(t) - y \left[1 + \mu^2 (x-y)^2\right]$$

Since
$$\alpha \equiv \frac{t_d}{t_r}$$
 is very large,
 $x = 1 + O(\alpha^{-1})$
 $y' = p(t) - y [1 + \mu^2(1 - y)^2] + O(\alpha^{-1})$

$$x \approx 1$$

$$y' \approx \frac{p(t)}{p(t)} - y \left[1 + \mu^2 (1 - y)^2\right]$$

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The box model: one state variable

$$y' = p(t) - y \left[1 + \mu^{2}(1 - y)^{2}\right]$$

= $\bar{p} + \hat{p}(t) - y \left[1 + \mu^{2}(1 - y)^{2}\right]$
$$\hat{p}(t) = 0 \implies y' = -\frac{\partial}{\partial y} \left[\mu^{2}(\frac{y^{4}}{4} - \frac{2y^{3}}{3} + \frac{y^{2}}{2}) + \frac{y^{2}}{2} - \bar{p}y \right]$$



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$$y' = \overline{p} + \hat{p}(t) - y \left[1 + \mu^2 (1 - y)^2\right]$$

Let $\hat{p}(t) = \begin{cases} 0 & t \le 0 \\ \Delta & 0 \le t \le \tau \\ 0 & t > \tau \end{cases}$



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Deterministic perturbations



FIG. 4. The critical perturbation, of amplitude Δ_0 , is found by requiring that the perturbed potential, shown here as a function of y, has one minimum and one inflection point. The perturbed potential is obtained by replacing \bar{p} with $\bar{p} + \Delta$ in (2.9) and requiring that the global minimum y_e of Fig. 2 becomes the inflection point y'_a .

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For $\Delta > \Delta_0$, how long can the system tolerate $p = \bar{p} + \Delta$ and still recover to y_a ?

$$\frac{dy}{dt} = \overline{p} + \Delta - y \left[1 + \mu^2 (1 - y)^2\right]$$
$$\frac{dy}{\overline{p} + \Delta - y \left[1 + \mu^2 (1 - y)^2\right]} = dt$$
$$\int_{y_a}^{y_b} \frac{dy}{\overline{p} + \Delta - y \left[1 + \mu^2 (1 - y)^2\right]} = \int_0^\tau dt = \tau$$

[Why does this work?]



Deterministic perturbations



FIG. 3. The minimum amplitude of a perturbation, as a function of its duration, that will shift the system from the globally stable equilibrium y_a of Fig. 2 to the metastable state, y_c . The perturbation must exceed a critical amplitude, Δ_0 , in order to displace the system from y_a , even if applied for an infinite time.

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Cessi's calculation

- 1,000 years $\longrightarrow \tau = 4.6$
- Critical value of Δ for au= 4.6 is $\Delta\approx$ 0.3
- $\Delta_0 = 0.2$ corresponds to freshwater flux of 0.4 $m yr^{-1}$
- Max meltwater flux preceeding Younger Dryas was 0.5 $m yr^{-1}$

"Close"

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Questions to pursue

- What about $\hat{p}(t)$ continuous?
- How do $\mathcal{O}(\alpha^{-1})$ terms affect results?
- Extend calculations to higher dimensional systems?

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Reference

Cessi, Paola (1994). A simple box model of stochastically forced thermohaline flow. *Journal of Physical Oceanography* v.24, pp. 1911-1920.